

Metabolic strategies of free-living and aggregate-associated bacterial communities inferred from biologic and chemical profiles in the Black Sea suboxic zone

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Abstract

The Black Sea is a permanently anoxic basin with a well-defined redox gradient. We combine environmental 16S rRNA gene data from clone libraries, terminal restriction fragment length polymorphisms, and V6 hypervariable region pyrosequences to provide the most detailed bacterial survey to date. Furthermore, this data set is informed by comprehensive geochemical data; using this combination of information, we put forward testable hypotheses regarding possible metabolisms of uncultured bacteria from the Black Sea's suboxic zone (microaerophily, nitrate reduction, manganese cycling, and oxidation of methane, ammonium, and sulfide). Dominant bacteria in the upper suboxic zone included members of the SAR11, SAR324, and Microthrix groups and in the deep suboxic zone included members of BS-GSO-2, Marine Group A, and SUP05. A particulate fraction (30 µm filter) was used to distinguish between free-living and aggregate-attached communities in the suboxic zone. The particulate fraction contained greater diversity of V6 tag sequences than the bulk water samples. Lentisphaera, Epsilonproteobacteria, WS3, Planctomycetes, and Deltaproteobacteria were enriched in the particulate fraction, whereas SAR11 relatives dominated the free-living fraction. On the basis of the bacterial assemblages and simple modeling, we find that in suboxic waters, the interior of sinking aggregates potentially support manganese reduction, sulfate reduction, and sulfur oxidation.

Introduction

The Black Sea, a semi-enclosed basin, contains at least three distinct microbial ecosystems. The surface layer of the Black Sea is well-oxygenated and driven by oxygenic photosynthetic processes, whereas the aphotic deep layer is anoxic and sulfidic. A suboxic zone (c. 50 m thick) lies at the boundary between the oxic and anoxic layers. The suboxic zone, which has < 10 µM oxygen and undetectable hydrogen sulfide (Murray *et al.*, 1995) hosts a variety of microbial metabolisms including anoxygenic photosynthesis, manganese oxidation, nitrification, denitrification, and anammox (Kuypers *et al.*, 2003; Manske *et al.*, 2005; Lam *et al.*, 2007; Oakley *et al.*, 2007; Clement *et al.*, 2009). The hydrological balance is influenced by freshwater input from rivers such as the Danube, which mix into surface waters, whereas salty Mediterranean waters flow through the Bosporus Straight and fill the deep basin. These two different water sources cause the Black Sea to be permanently stratified with respect to salinity and density, and this has led to physical, chemical (Caspers, 1957; Sorokin, 1983), and therefore microbial stratification of the Black Sea's water column.

The Black Sea is an ideal place to study microbially mediated redox reactions in low oxygen conditions. The vertical scale of the highly stable redox gradient is in the order of meters to tens of meters, providing the opportunity to finely sample the sequence of redox reactions (Murray *et al.*, 1995). Nonetheless, bacterial communities of the Black Sea suboxic zone have not been thoroughly studied. Other researchers have examined bacterial functional genes (Lam *et al.*, 2007; Oakley *et al.*, 2007), specific bacterial groups in the suboxic layer (Manske *et al.*, 2005; Kirkpatrick *et al.*, 2006; Schubert *et al.*, 2006), or bacteria in the sulfidic zone (e.g. Grote *et al.*, 2008; Glaubitz *et al.*, 2010). Lin *et al.* (2006) quantified some phylum-level groups in the suboxic zone using FISH, but only Vetriani *et al.* (2003) examined specieslevel diversity of the general bacterial community.

Due to the stratification of the Black Sea, characteristic inflections in the water-column profiles of nitrate, manganese, cesium isotopes, and mesoplankton, etc. are associated with specific density values, but not with specific depths, regardless of when and where they were sampled (Vinogradov & Nalbandov, 1990; Buesseler *et al.*, 1991; Codispoti *et al.*, 1991; Lewis & Landing, 1991; Murray *et al.*, 1995). Therefore, results presented here will be plotted against potential density (σ_0) rather than depth (m). Both axes are used in Fig. 1 for comparison.

In most of the Black Sea, the oxygenated cold intermediate layer (core density of $\sigma_0 = 14.5$) represents the lower boundary of direct communication with the surface; the ventilation frequency of this layer depends on winter conditions (Tolmazin, 1985; Gregg & Yakushev, 2005). However, the suboxic layer is still linked to euphotic processes via the sinking of organic matter. In marine systems, organic matter aggregates have been found to be enriched in nutrients and to exhibit higher bacterial abundance and enzyme activity compared with the water



Fig. 1. Concentration of suspended particulate organic carbon. Both potential density and depth axes are shown for comparison. Dotted lines indicate the borders of the suboxic zone. Dashed line indicates the cold intermediate layer (CIL).

column (Simon *et al.*, 2002). The bacterial communities of free-living and aggregate-attached bacteria have been reported to be taxonomically distinct in marine environments (e.g. DeLong *et al.*, 1993; Kellogg & Deming, 2009), but aggregate-attached bacterial communities in suboxic waters have not been previously examined.

In this article, we report the most comprehensive census of microbial diversity for the suboxic Black Sea published to date, and we examine the contribution of sinking aggregates to the taxonomic composition of the suboxic zone. Many bacteria in the Black Sea are not closely related to cultured organisms (Vetriani et al., 2003) and have unknown metabolisms. Therefore, in addition to describing our census of diversity, we attempt to make testable hypotheses regarding the likely metabolisms of uncultured organisms by linking our biological data to a comprehensive suite of chemical measurements. Our biological data includes three complementary and independent datasets: full-length16S rRNA gene clone libraries, V6 hypervariable region tag pyrosequences, and terminal restriction fragment length polymorphism (TRFLP) profiles. The large number of V6 tags (c. 9000 per sample) allows deep examination of diversity and community similarities, whereas TRFLP analyses of many samples yield a detailed depth profile. Our integration of multiple biologic and chemical datasets represents the most thorough description of Black Sea microbial ecology to date.

Materials and methods

Sampling

DNA samples were collected in the western central gyre of the Black Sea in late March 2005 on cruise 403 of the *R/V Endeavor* (42°30' N, 30°45' E). Samples were collected using a CTD-Rosette with 10-L Niskin bottles and Sea Bird sensors. Approximately 2 L of sample were filtered into 0.2- μ m Millipore Sterivex filters for bulk water. Aggregate-associated samples were collected using a 30- μ m flat polypropylene filter (Millipore) from a separate cast at the same station. Samples were immediately frozen and stored at -80 °C upon arrival in the laboratory. Some full-length sequences were also obtained from samples collected in the same manner in April 2003 on Voyage 162 leg 17 of the *R/V Knorr* (42°30' N, 31°00' E).

DNA extraction

The DNA extraction protocol was adapted from Vetriani *et al.* (2003) and included 8–10 freeze thaw cycles between a dry ice/ethanol bath and a 55 $^{\circ}$ C water bath followed by chemical lysis with lysozyme and proteinase

K. All DNA used for V6 tag amplification and TRFLP was extracted in this fashion. The 30 μ m filter used for 16S rRNA gene clone libraries was extracted with an additional bead beating step utilizing 0.1 and 0.5 mm zirconia-silica beads.

Amplification of 16S rRNA gene V6 hypervariable region

Four samples from March 2005 were amplified targeting the V6 hypervariable region of 16S rRNA gene: bulk water from densities $\sigma_{\theta} = 15.34$, 15.79, 16.08, and a 30 µm filter from $\sigma_{\theta} = 15.8$. V6 amplicon libraries were constructed and sequenced as in Huber *et al.* (2007). PCR amplification, in 30 cycles, used 0.2 µM of each primer in a cocktail of five forward primers (967F) and four reverse primers (1046R; as per Huber *et al.*, 2007). Multiple primers were used to increase taxonomic coverage. Eighty-nine percent of all sequences in the Greengenes database (DeSantis *et al.*, 2006; accessed April 2009) matched the V6 primer set, although 18% of *Planctomycetes* sequences had a mismatch.

PCR products were pyrosequenced on a Roche Genome Sequencer FLX in the Sogin lab (Marine Biological Laboratory, Woods Hole, MA) as part of the International Census of Marine Microbes (ICoMM). The resulting tag sequences were screened for quality as recommended by Huse *et al.* (2007). Tag sequences have been deposited in the National Center for Biotechnology Information (NCBI) Short Read Archive (SRA) under the accession numbers SRS005799–SRS005802.

Although we could not obtain replicate pyrosequencing samples, Brazelton *et al.* (2010) obtained two replicate pyrosequencing runs from the same facility. These replicate sequencing runs were performed months apart with different amplicon libraries. The Bray–Curtis similarity of these replicates was 89%, greater than any of the similarity values reported in this study.

Alignment of V6 tags

Bacterial sequence alignments were constructed by submitting the unique V6 tag sequences from all four samples to the NAST aligner (http://greengenes.lbl.gov). Primers were included to ensure full-length alignment. The alignment was manually corrected and primers were trimmed. The distance matrix for each alignment was calculated with quickdist as described in Sogin *et al.* (2006) without penalties for terminal gaps. Sequences were clustered into operational taxonomic units (OTUs), and rarefaction curves and diversity estimators were calculated with DOTUR (Schloss & Handelsman, 2005).

Comparisons of V6 tags between samples

The program sons (Schloss & Handelsman, 2006) was used to determine the relative abundance distribution of each OTU in each sample. To normalize the abundances of each sequence type among samples, tags were randomly resampled using Daisychopper (available at http:// www.genomics.ceh.ac.uk/GeneSwytch/Tools.html). These normalized OTUs were used in Venn diagrams, depth profiles, and Bray–Curtis and Jaccard similarity calculations (calculated with PRIMER 6 (http://www.primer-e. com) without any further data transformation). Bray– Curtis similarities will be shown here as the index utilizes abundance data. Jaccard similarities (presence/absence) are lower, but follow the same trend.

16S rRNA gene clones

Nearly full-length 16S rRNA gene clones were amplified from suboxic water collected from the western central gyre in April 2003, and from $\sigma_{\theta} = 15.3$ and the 30-µm filter from $\sigma_{\theta} = 15.8$ in the western central gyre in March 2005. PCR was performed using universal bacterial primers 27F (5'-AGAGTTTGATCCTGGCTCAG-3') and 1492R (5'-GGTTACCTTGTTACGACTT-3') for 32 cycles with annealing at 60 °C (Lane, 1991). Sequencing was done at High-Throughput Sequencing Solutions (http://www. htseq.org) using primers T7, M13R, and 357F (Muyzer *et al.*, 1993). Chromatograms were hand-inspected and contigs assembled using the Sequencher program (Gene-Codes Corporation, Ann Arbor, MI). SSU rRNA gene sequences were deposited at Genbank under the accession numbers GU145386–GU145550.

Taxonomy

Taxonomy was assigned to each V6 tag sequence by the GAST process (Huse et al., 2008). Tags with distances > 0.25 were added to the unclassified category. V6 tags were also compared with Black Sea full-length 16S rRNA gene sequences. The V6 regions of the full-length sequences were compared with the V6 tags using MATGAT 2.0 (Campanella et al., 2003) to obtain distances. The taxonomy of nearly full-length Black Sea sequences was determined by the Greengenes classifier using Hugenholtz taxonomy (http://greengenes.lbl.gov). Boot-strapped neighbor-joining trees were created in Arb after alignment to a master database using the NAST tool of greengenes.lbl.gov. To clarify differences between Marine Group A and Deferribacteres, a tree was originally created using sequences used to define the Deferribacteres and Marine Group A phyla in Jumas-Bilak et al. (2009). To save room, many Deferribacteres sequences were removed in the final figure.

TRFLP

TRFLPs were obtained from a profile (14 depths) of the western central gyre in March 2005 using universal bacterial primers 27F-FAM and 1517R (5'-ACGGCTACC TTGTTACGACTT-3') (Vetriani et al., 2003). PCR products were amplified for 30 cycles at 48 °C using 2 × PCR MasterMix (Fermentas, Ontario, Canada). Purified PCR products (QiaQuick columns; Qiagen, Valencia, CA) were separately digested overnight with four restriction enzymes (HaeIII, Hpy1881, MspI, MnII) and immediately ethanol precipitated according to the manufacturer's instructions (Amersham Pharmacia Dynamics). Analysis was performed on a MegaBACE 1000 apparatus (Molecular Dynamics) at the University of Washington Marine Molecular Biotechnology Laboratory. Electrophoretic profiles were visualized using Dax software (Van Mierlo Software Consultancy, The Netherlands). TRFLP profiles were normalized by total peak height. TRFLP peaks were binned using frame shifting (Hewson & Fuhrman, 2006) with four frames at 0.5-bp intervals. For each enzyme, a resemblance matrix was obtained using either the Bray-Curtis index, which takes abundance (peak height) into account, or the Jaccard index which uses presence/ absence, and for each comparison between two samples, the maximum similarity of the four frames was used. Profiles were clustered using the PRIMER 6 program. Error in the resemblance matrix and significance level of the cluster diagram was determined with a Monte-Carlo simulation of 50 replicates using the average standard deviation in both peak height and base pairs as determined by 16 pairs of replicate TRFLP profiles.

Select 16S rRNA gene clone PCR products from 2003 and 2005 were digested with all four restriction enzymes and used to identify TRFLP peaks with a range \pm 0.5 bp from the length of the digested clone. For a peak to be considered positively identified as a clone library sequence via TRFLP, corresponding peaks must have been present in electrophoretic profiles produced by two or more endonucleases, and a match must have been made for the shape of peak height profile vs. depth for two or more enzymes.

Due to the replicability of the relative peak heights and the lack of cloning bias (Rainey *et al.*, 1994), and because each PCR was run under the same conditions with similar extracts from the same amount of material, each normalized TRFLP peak was compared between different TRFLP profiles. However, due to PCR bias (Polz & Cavanaugh, 1998), comparison between the heights of different peaks was made with caution. More than one bacterial species can produce the same TRFLP peak; however, by ensuring that the shape of a peak's depth profile must match between more than one enzyme, and using a small bin size, that risk was reduced.

Cell counts

Ten milliliters of water from density surfaces $\sigma_{\theta} = 15.5$, 15.7, 15.9, and 16.0 was filtered onto a 0.2 μ m filter (Millipore) and frozen. DAPI staining was performed in the laboratory, and a minimum of 200 cells were counted.

Chemical data

Oxygen was measured using the classic Winkler method, and sulfide was measured by iodometric titration (Cline, 1969). In both cases, reagents were bubbled with argon to avoid contamination by atmospheric oxygen. Nutrients were analyzed using a two channel Technicon Autoanalyzer II system. Nitrate was reduced to nitrite using a cadmium column, which was measured using sulfanilamide and N(1-naphthyl)-ethylenediamine (Armstrong et al., 1967). Ammonium was analyzed using the indophenole blue procedure (Slawky & MacIsaac, 1972). Deep water samples were diluted with nitrate-free Black Sea surface water to reduce sulfide content. Particulate manganese was filtered onto 0.4 µm filters. Oxidized particulate manganese was determined by B. Tebo by reducing the particulate Mn in 0.1% hydroxylamine and then measuring Mn(II) using the formaldoxime method (Brewer & Spencer, 1971) as seen in Konovalov et al. (2003).

Methane was measured in water samples collected for N₂/Ar in evacuated 250-mL glass flasks. In the half-full flasks, the water was equilibrated with the headspace overnight and then removed. Gas samples were cryogenically processed and measured at the Stable Isotope Lab, School of Oceanography, University of Washington, on a Finnegan Delta XL isotope ratio mass spectrometer using interfering masses for mass 32 and 16. Millivolts from mass 32 multiplied by the ratio of 16/32 measured for air was subtracted from millivolts of mass 16. The remainder was considered to be mV of methane. Methane data obtained from the sulfidic layer of the Western Central Gyre in 2003 using gas chromatograph methods (A.V. Egorov, unpublished data) was used to calibrate mV of methane $(R^2 = 0.99)$ with the assumption that methane concentration at the same density in the sulfidic zone are relatively constant over time, which is consistent with the data of Kessler et al. (2006).

V6 tag sequence depth profiles

V6 tag sequences were categorized into five different depth profiles based on the normalized number of tags at each depth using a Perl script. In each category, at least one depth must have had 10 or more tags of that particular sequence, to be sure that enough information existed to categorize the profile properly. Due to the random nature of the normalization, some OTUs with a low frequency of tags (e.g. Chlorobi phylotype BS130 with 43 original but < 10 normalized tags at $\sigma_{\theta} = 16.1$) were not categorized, but might become categorized if the normalization were repeated. The five potential metabolic group profile categories (Metabolic Groups I–V) were defined by the following equations. The number of tags of a particular sequence at the density indicated inside the brackets, for example {15.3}, equals the number of tags of a particular sequence at the density $\sigma_{\theta} = 15.3$.

Metabolic Group I (aerobic; Fig. 6a): $\{15.3\} \ge 10$ tags $\cap 1.55 \times \{15.8\} < \{15.3\} \cap \{16.1\} < 10$ tags.

Metabolic Group II (nitrate reduction; Fig. 6b): $\{15.3\} \ge 10 \text{ tags } \cap 1.55 \times \{15.3\} > \{15.8\} > 0.45 \times \{15.3\} \cap \{16.1\} \text{ is either } < 10 \text{ tags } \cup < 0.03 \times \{15.3\}.$

Metabolic Group III (manganese oxidation; Fig. 6c): $\{15.8\} > 1.55 \times \{15.3\}$ and $> 1.55 \times \{16.1\} \cap \{15.8\} \ge 10$ tags.

Metabolic Group IV (methane oxidation or manganese oxide reduction; Fig. 6d): $\{16.1\} \ge 10$ tags and $\ge \{15.8\} \cap \{15.8\} \ge 5$ tags $\cap \{15.3\} < 5$ tags.

Metabolic Group V (sulfur cycling; Fig. 6e): {15.3} and {15.8} < 5 tags \cap {16.1} \geq 10 tags.

Tags not fulfilling the requirements for any of these categories were labeled as uncategorized. 'Aggregate-attached' and 'free-living' designations were also determined with a Perl script (Fig. 5c).

Aggregate-attached only: [30 μm] \geq 4 \times [15.8] \cap [30 μm] \geq 10 tags.

Free-living only: [15.8] \geq 5 \times [30 μm] \cap [15.8] \geq 10 tags.

Aggregate-associated and free-living (i.e. abundant on both filters): $[30 \ \mu\text{m}] < 4 \times [15.8] \cap [15.8] < 5 \times [30 \ \mu\text{m}] \cap [30 \ \mu\text{m}] \ge 10 \text{ tags.}$

Cut-offs for aggregate-attached ($4\times$ enriched) and freeliving ($5\times$ enriched) were determined empirically. The classification of OTUs as free-living, versatile, or aggregate-attached did not depend greatly on the threshold criteria. If free-living bacteria were determined by a threshold of $3\times$ more abundance in the bulk water instead of $4\times$, then three more OTUs would become designated as free-living. If the threshold for aggregateattached bacteria were changed to $3\times$ more abundance in the particulate sample, then four OTUs would become aggregate-attached.

Results

Chemical profiles

Chemical concentrations and their fluxes at the depths sampled for V6 pyrosequencing are seen in Table 1. Full depth profiles are shown in Fig. S1. Oxygen decreased



	$\sigma_{\theta} = 15.3$		$\sigma_{\theta} = 15.8$		$\sigma_{\theta} = 16.1$	
	Flux	Conc.	Flux	Conc.	Flux	Conc.
0 ₂	13 251	39.5	2141	1.8		b.d.
NO_3^-	22*	4.8	208	2.1		b.d.
NH_4^+		b.d.	127	0.006	312	2
CH_4		b.d.	14	b.d.	34	0.2
H₂S		b.d.		b.d.	107	b.d.
PMn	No gradient	0.016	920	0.054	2.8	0.031

b.d., below detection. Fluxes were calculated using mixing coefficients from Ivanov & Samodurov (2001).

*At the nitrate maximum, diffusion was a negative flux.

from 335 μ M at the surface ($\sigma_{\theta} = 14.29$) to 10 μ M at $\sigma_{\theta} = 15.65$, and was undetectable below $\sigma_{\theta} = 16.0$ (σ_{θ} = 17.21). The first detectable sulfide (3 μ M; Konovalov et al., 2003) was at $\sigma_{\theta} = 16.11$, which was slightly deeper than the deepest V6 sample ($\sigma_{\theta} = 16.08$). Sulfide then increased to 380 µM in the deep water. Nitrate was below 0.1 µM for the top 58 m of the Black Sea $(\sigma_{\theta} = 14.47)$ but then increased to a maximum of 4.8 μ M at $\sigma_{\theta} = 15.38$ and then decreased to undetectable at $\sigma_{\theta} = 15.9$. Nitrite had a maximum of 0.09 μ M at $\sigma_{\theta} = 15.85$. Ammonium concentrations were 0.08 μ M at $\sigma_{\theta} = 15.85$ and increased to 2.9 μ M at the bottom of the suboxic zone and then to 98 µM at depth. Methane was first measurable at $\sigma_{\theta} = 15.85$, and then increased to 0.4 μ M at the bottom of the suboxic zone and to 13.5 μ M at 750 m (σ_{θ} = 17.15). The concentration of particulate manganese was variable, but low in the oxycline. Particulate manganese had a maximum of 0.22 µM at $\sigma_{\theta} = 15.85$ and decreased to *c*. 0.01 μ M in deeper water.

Suspended particulate carbon (S-POC) concentrations are shown in Fig. 1. S-POC was 10.2 μ M at 20 m in the euphotic zone, and then decreased to between 5 and 7 μ M from 50 to 57 m ($\sigma_{\theta} = 14.4-14.6$). S-POC increased to 13.7 μ M at $\sigma_{\theta} = 15.1$, and then decreased to 4 μ M in the suboxic zone.

Microbial data

TRFLP chromatograms were obtained from 14 depths including nine in the hypoxic/suboxic region. These TRFLP chromatograms (MspI) illustrate the bacterial community shifts with depth (Figs S2–S4). Comparison of TRFLP profiles (Fig. S2) indicate that all the low oxygen bacterial communities were significantly different (< 25% similarity) both from the community at higher oxygen concentrations and from the community in the deep sulfidic zone ($\sigma_{\theta} = 16.8$ [275 m], 17.19 [1000 m]

and 17.21 [2000 m]). However, $\sigma_{\theta} = 16.1$ shared some of the bacterial community (41% similarity) with the sulfidic sample $\sigma_{\theta} = 16.4$ (141 m). Interestingly, samples from the suboxic zone formed three separate bacterial communities within a larger coherent cluster (Fig. S2): upper suboxic ($\sigma_{\theta} = 15.3-15.7$), lower suboxic ($\sigma_{\theta} = 15.75-$ 15.95), and deep suboxic/upper sulfidic ($\sigma_{\theta} = 16.0-16.4$). The water samples used in pyrosequencing represent these three distinct bacterial communities, and span varying oxygen and nutrient concentrations (Table 1; Fig. S1).

We obtained 36 342 high quality bacterial V6 tag sequences from four samples, representing a total of 2088 OTUs at 0.03 distance (97% similarity) with a range of 726–888 OTUs and 7761–10 566 tags per sample. We identified 18 of the OTUs in TRFLP profiles (30% of TRFLP peaks) by linking V6 tag sequences to corresponding full-length 16S rRNA gene clones. The taxonomic composition of each sample, based on V6 pyrosequences, is

shown in Fig. 2 (first column). The most frequently occurring V6 tag sequences in each sample are shown in the second column of Fig. 2. The most dominant sequence in the $\sigma_{\theta} = 15.3$ and 15.8 samples matched the full-length sequence BS007 from group II of the SAR11 clade of Alphaproteobacteria (Fig. 3), which are typically oligotrophic heterotrophs found in the mesopelagic ocean (Carlson et al., 2009). An uncultured member of the Gammaproteobacteria, which matched full-length sequence BS129 from group BS-GSO2 (Fig. 3), was the most dominant V6 sequence at $\sigma_{\theta} = 16.1$. BS129 was quite closely related to II8-19 (GU108534) (Fig. 3), which was found to be autotrophic during stable isotope probing of the upper sulfidic zone of the Black Sea (Glaubitz et al., 2010). The SAR324 group of Deltaproteobacteria was abundant in every sample. However, the SAR324 V6 tag sequence abundant at $\sigma_{\theta} = 15.3$ could not be linked to a full-length clone sequence. V6 tags matching BS134 dominated at $\sigma_{\theta} = 15.8$. V6 tags matching



Fig. 2. Taxonomic composition of (a) $\sigma_{\theta} = 15.3$, (b) $\sigma_{\theta} = 15.8$, (c) $\sigma_{\theta} = 16.1$, (d) > 30 µm fraction from $\sigma_{\theta} = 15.8$ as determined by the V6 hypervariable region of 16 S rRNA. The first column contains pie charts of all taxonomic groups with 20 or more tags. The second column contains OTUs that are > 1% of total tags. Bar graphs and pie charts share the same legend. Bars are labeled by 100% matches to full-length sequences (BSXXX) and more specific taxonomic information. Some abundant V6 tags had 100% matches to a sequence not from the Black Sea. In these cases, the accession number is shown. BS158 and BS079, Actinobacteria related to the deeply branching microaerophilic heterotroph *Microthrix parvicella* (Rossetti *et al.*, 2005; Fig. 4), were also present in all samples. V6 tags matching BS110, a member of Marine Group A (Fig. 4), dominated the aggregate sample along with tags matching BS134 (SAR324), BS158 (*Microthrix*), and BS109, an unclassified *Planctomycetes* (Fig. 4). BS109 primarily groups with other Black Sea sequences sequenced using *Planctomycetes* specific primer sets (Fig. 4; Kirkpatrick *et al.*, 2006; Woebken *et al.*, 2008).

The GAST process, which utilizes SILVA taxonomy (Huse *et al.*, 2008), assigned many tag sequences as *Deferribacteres*. However, all '*Deferribacteres*' sequences from the GAST database with matches to Black Sea tags appear to belong to Marine Group A (Fig. 4). Pyrosequences from this uncultured phylum (BS110, BS100, BS137, BS042) are found in all four samples.

While all V6 tag samples were obtained from depths containing low oxygen and no measurable sulfide, the

39 μ M oxygen (σ_{θ} = 15.3) (Fig. 2a) and the undetectable oxygen ($\sigma_{\theta} = 16.1$) (Fig. 2c) samples had dissimilar bacterial communities (11% Bray-Curtis similarity), whereas the 39 μ M oxygen (σ_{θ} = 15.3) and 2 μ M oxygen samples $(\sigma_{\theta} = 15.8)$ (Fig. 2b) had more similar communities (51% Bray–Curtis). The 30 μ M oxygen ($\sigma_{\theta} = 15.3$) and 2 μ M oxygen (σ_{θ} = 15.8) samples also share more OTUs (248) than do the 2 µM oxygen and the deep suboxic zone samples ($\sigma_{\theta} = 16.1$) (154) (Fig. 5b). The bacterial community on the 30 µm filter was most similar (43%) to the community from bulk water from the same depth $(\sigma_{\theta} = 15.8)$, but was more similar to the sample from $\sigma_{\theta} = 16.1$ (33%) than to $\sigma_{\theta} = 15.3$ (24%). The particulate sample shared significant OTUs with all three depths (Fig. 5b), but shared many OTUs (111), otherwise only found at $\sigma_{\theta} = 16.1$.

The diversity between samples varied greatly. The undetectable oxygen sample ($\sigma_{\theta} = 16.1$) and particulate sample ($\sigma_{\theta} = 15.8$) showed significantly more diversity



Fig. 3. A neighbor-joining phylogenetic tree of proteobacterial groups important in the Black Sea suboxic zone. Black Sea sequences discussed in the text are in red. Other Black Sea sequences (from this work or from Vetriani *et al.*, 2003 or Glaubitz *et al.*, 2010) are shown in pink. Cultured, enriched, or sequenced organisms are in blue. Asterixes indicate autotrophic bacteria identified using stable isotope probing with bicarbonate (Glaubitz *et al.*, 2010). Outgroup is *Flexibacter litoralis* (M58784).



Fig. 4. A neighbor-joining phylogenetic tree of nonproteobacterial groups important in the Black Sea suboxic zone. Black Sea sequences discussed in the text are in red. Other Black Sea sequences (Vetriani *et al.*, 2003; Kirkpatrick *et al.*, 2006; Woebken *et al.*, 2008; Glaubitz *et al.*, 2010) are shown in pink. Cultured, enriched or sequenced organisms are in blue. This tree indicates that Marine Group A, not *Deferribacteres*, is an important phylum in the Black Sea. The tree was created using sequences used to define the Marine Group A phyla in Jumas-Bilak *et al.* (2009). Sequences with names starting with V6 were used in the GAST process (Huse *et al.*, 2008) to assign Black Sea V6 tags. Outgroup is *Thermotoga* sp str KOL6.

than the bulk water samples with low (30 μ M; $\sigma_{\theta} = 15.3$) or very low (2 μ M oxygen; $\sigma_{\theta} = 15.8$) oxygen levels, as seen in rarefaction curves (Fig. 5a). Chaol indices for the samples range from 900 to 1250, indicating significantly less diversity than seen in Atlantic Ocean seawater with similar sequencing effort (Chaol = 13 772) (Sogin *et al.*, 2006). Even when all Black Sea samples were pooled, they showed less diversity (Chaol = 3529).

Comparison between methods

Pyrosequencing and TRFLP both avoid cloning biases (Rainey *et al.*, 1994), but still contain PCR biases (Polz & Cavanaugh, 1998; Huse *et al.*, 2007, 2008). Despite the use of different primers, conclusions from TRFLP data and V6 tag sequences compare well. We can identify many of the same OTUs using both techniques. Not only are depth profiles of individual OTUs similar between



Fig. 5. Comparison of samples from which V6 tag sequences were obtained. (a) Rarefaction curves of OTUs at 0.03 distance for the four samples with V6 tags and a pooled sample combining all four samples. The pooled sample includes 36 342 tags and 2088 OTUs, but the entire curve is not shown for clarity. (b) Venn diagram of OTUs at 0.03 distance shared by the four samples. The small ovals indicate that four OTUs were shared only between $\sigma_{\theta} = 15.3$ and $\sigma_{\theta} = 16.1$, whereas 66 OTUs were shared only between $\sigma_{\theta} = 15.8$ and the aggregate-associated sample. (c) The broad taxonomic designations of aggregate-attached (5× more abundant in the 30 µm filter), free-living (4× more abundant in 0.2 µm filter), and bacteria abundant in both fractions from density $\sigma_{\theta} = 15.8$. The outer ring indicates the relative abundance of each OTU. Taxonomic groups with < 30 tags were not included in (c) for simplicity, but can be found in Tables S1–S8.

techniques (Fig. 6) but similarity indices are also similar. The Bray–Curtis similarity between $\sigma_{\theta} = 15.3$ and 15.8 is 52% for TRFLP and 50% for V6 tags with > 1% relative abundance. The Bray–Curtis similarity between $\sigma_{\theta} = 15.8$ and 16.1 is 40% and 43%, respectively.

Our V6 pyrosequence dataset corresponds well with previously published quantitative analyses. For these comparisons, it is important to note that cell counts from the suboxic zone at this station (5 ± 1E5 cells mL⁻¹; Table S1) were similar to values obtained on other cruises (Lin *et al.*, 2006), and that cell counts did not change appreciably throughout the suboxic zone (Table S1). Furthermore, the percentage of total V6 tag pyrosequences at each density surface was similar to previously reported values determined using more quantitative methods. For example, the proportion of V6 tag sequences at $\sigma_0 = 16.1$

assigned to the family Methylcoccales (1%) is identical to the proportion of bacteria identified as Methylococcales (1%) in 2001 using quantitative PCR of 16S rRNA gene (Schubert et al., 2006). The proportion of cells with bacterial chlorophyll e (0.5-1%) in 2001 (Manske et al., 2005) is similar to the proportion of V6 tag sequences assigned to Chlorobi (0.6%). The proportions of Candidatus Scalindua sorokinii in 2001 (0.8%; Kuypers et al., 2003) and sulfate reducers in 2003 (8%; Lin et al., 2006) identified using FISH also matched closely with the proportion of V6 tag sequences assigned to these groups (0.8% and 11%, respectively). However, a few discrepancies should be noted. Epsilonproteobacteria were 1.4% of V6 tags at σ_{θ} = 16.1 in 2005, but 6% of DAPI stained cells were found to be Epsilonproteobacteria by FISH at the same station in 2003 (Lin et al., 2006).

Gammaproteobacteria comprise 25% of V6 tag sequences, but only 6% were found using FISH (Lin *et al.*, 2006). As the *Gammaproteobacteria* are a phylogenetically diverse group, that discrepancy may be due to FISH probe mismatches to some sub-groups. Most surprisingly, *Bacteroidetes* comprised only 0.4% of V6 tag sequences, but 5% of cells were identified as *Bacteroidetes* in 2003 using FISH (Lin *et al.*, 2006). It is puzzling that TRFLP fragments corresponding to *Bacteroidetes* sequences BS040 and BS035 (both appear as a 91-bp fragment using the MspI enzyme in Fig. S3) were identified in $\sigma_{\theta} = 15.8$, but no corresponding V6 tags were found. No mismatch was found between these sequences and the V6 primer set. In summary, there is general agreement between the proportion of V6 pyrosequences and published quantitative analyses.

(a) Metabolic group I O2 (µM) (aerobic) 150 50 100 14.8 TRFLP V6 tags 15.2 Sigma theta 15.0 Arctic96BD-19 BS028 o O. 16.4 76 OTUs (4192 tags) 100 150 200 250 0 50 (b) Metabolic group II NO3- (µM) 2 3 (NO,- red.) 14.8 0 0 15.2 Sigma theta 9 08 15.0 16.0 SAR11 BS007 ¢ NO 16 21 OTUs (4140 tags) 800 1600 2400 (c) Metabolic group III Particulate Mn (µM) 0 0.05 0.1 0.15 0.2 0.25 (Mn oxid.) 14.8 SAR324 BS134 MnO. 15.3 ma theta 15.0 Sig 16.0 16.4 32 OTUs (4959 tags) 500 1000 1500 (d) Metabolic group IV NH4+ and CH4 (µM) Particulate Mn (µM) (CH, oxid. / Mn red.) 4 6 8 10 n 0.05 0.1 0.15 0.2 0.25 14.8 mma BS129 Marine group A BS137 _____MnO, NH. + CH 15.2 Sigma theta 15.0 16.0 Δ 16 22 OTUs (3372 tags) 400 800 1200 1600 2000 400 800 1200 Normalized V6 tags (e) Metabolic group V H,S (µM) 10 20 30 (S cycling) 14.5 Sulfurimonas BS139 VH,S 15.2 Sigma theta 15.0 16.0 16. 50 100 150 200 53 OTU (1178 tags) Normalized V6 tags

Fig. 6. Depth profiles of bacteria in the Black Sea and their taxonomy. Five bacterial depth profiles were predicted from chemical profiles and fluxes (Table 1). Pie charts indicate the taxonomy of all the V6 tag sequences placed in each category, using the color key from Fig. 2. More specific taxonomies can be found in the Tables S1–S8. Outer ring indicates the number of OTUs in each group, and the size of each section in the outer ring indicates the number of tags in each OTU. An example fulllength sequence identified in both TRFLP chromatograms (in red) and V6 tag sequences (in blue) illustrates each depth profile. Relative fluorescent units for TRFLP peak heights are adjusted to match the scale for V6 tags. Profile types are as follows: (a) aerophilic, (b) nitrate reducing, (c) manganese oxidizing, (d) methane oxidizing, ammonia oxidizing, or manganese reducing, and (e) sulfur utilizing.

Discussion

The Black Sea is a permanently stratified basin. Chemical profiles occur in a predictable sequence on density surfaces through the Black Sea (Murray et al., 1995), although decadal (Konovalov & Murray, 2001), interannual (Fuchsman et al., 2008), and seasonal variability (Yakushev et al., 2006) in the concentrations of chemical species have been observed in some instances. The Black Sea has been stratified for over 7000 years (Jones & Gagnon, 1994), which is in contrast with transiently oxygenated basins such as the Baltic Sea (Hannig et al., 2007) or fjords such as Saanich Inlet (Manning et al., 2010), where periodic oxygenation events affect microbial activity. Thus, one expects less temporal variability in the microbial community of the Black Sea than seen in these transiently oxygenated basins, but this has never been examined. We can compare our data to TRFLP data (with restriction enzyme MnlI) from three suboxic water samples collected during a Black Sea expedition in 1988 (Vetriani et al., 2003). We have much greater confidence in our peak identifications, because we used multiple restriction enzymes (Engebretson & Moyer, 2003) and obtained TRFLP chromatograms of the clones themselves. We have, however, compared all TRFLP fragment sizes listed by Vetriani et al. with our TRFLP profiles. Of the TRFLP peaks listed by Vetriani et al. as being from the oxic zone, 45% were shared by our σ_{θ} = 14.9 sample from 2005. Of the peaks listed as being from the suboxic zone, 60% were shared by at least one suboxic zone sample from 2005. Therefore, the microbial community appears to show remarkable continuity from 1988 to 2005, but some differences are apparent, and more research is needed to more closely examine temporal variability in the Black Sea. In general, however, the stability of Black Sea's suboxic zone, when compared with transiently anoxic basins where mixing and flushing occur, simplifies the conceptual task of understanding why certain organisms may be present at specific depths.

By examining nine suboxic/hypoxic depths with TRFLP, we saw that although bacterial communities from the suboxic zone do form a coherent cluster, the suboxic zone separates into three distinct microbial communities: upper, lower, and deep suboxic zones (Fig. S2). This data support objections that typical use of the term 'suboxic' is too broad to define microbial communities and geochemical processes (Canfield & Thamdrup, 2009). However, the term suboxic is still useful to define a depth range in the water column of interest here. Each sample for V6 tag pyrosequencing in this study was obtained from a distinct geochemical regime (Table 1). At $\sigma_{0} = 15.3$, representing the upper suboxic zone, O₂ (39 µM) and NO₃⁻ (4.8 µM) were high (Fig. 6a and b), and the concentration of particulate organic carbon was

relatively high (6 μ M; Fig. 1). At σ_{θ} = 15.8, representing the lower suboxic zone, O_2 (1.8 μ M) and NO_3^- (2.1 μ M) decreased in concentration, and CH₄ and NH₄⁺ were also detectable (Table 1). The concentrations of particulate manganese oxides were at their maximum at $\sigma_{\theta} = 15.8$ (Fig. 6c), and so we consider this depth to be part of the manganese oxidation zone. At $\sigma_{\theta} = 16.1$, representing the deep suboxic zone, both O₂ and NO₃⁻ were undetectable, but CH₄ (0.2 μ M) and NH₄⁺ (2 μ M) were present along with a H₂S flux from sulfidic zone. The concentration of particulate manganese oxides at $\sigma_{\theta} = 16.1$ is less than at $\sigma_{\theta} = 15.8$ and continues to decrease with depth (Fig. 6c), and so we consider $\sigma_{\theta} = 16.1$ to be part of the manganese reduction zone. Thus, $\sigma_{\theta} = 15.3$, 15.8, and 16.1 represent three distinct geochemical regimes and three distinct bacterial communities (Figs S1 and S2; Table 1).

Predicting metabolisms

By examining the abundance of a bacterial taxon across large changes in geochemical gradients, one can make testable hypotheses about the metabolism of that taxon. We used normalized V6 tag abundance and normalized TRFLP peak height to represent bacterial abundance in this study. Our V6 tag abundances correspond well with previously reported values obtained with accepted quantitative techniques such as FISH and qPCR (see Results). Although the presence of DNA does not indicate metabolic activity, large differences in these DNA depth profiles that correspond to changes in the geochemical profiles probably reflect distinct zones of metabolic activity. This is particularly true for the Black Sea, which has maintained geochemical profiles at similar density surfaces at least since the 1960s (Konovalov & Murray, 2001) and probably much longer.

We used chemical fluxes (Table 1) to predict five general metabolic depth profiles (I-V) of organisms utilizing each oxidant or reductant: (I) oxygen utilization, (II) nitrate reduction, (III) manganese oxidation, (IV) oxidation of methane and ammonium or manganese reduction, and (V) sulfate reduction or sulfide oxidation. We then categorized each V6 OTU into one of these depth profiles (Metabolic Groups I-V). Known bacterial species falling into each of these depth profiles use metabolisms, consistent with the predictions (n = 46). In general, bacterial depth profiles measured by V6 tag sequences and by TRFLP peaks are equivalent (Fig. 6). Black Sea Groups I-V are quite different with respect to bacterial taxonomy (Fig. 6), and we discuss each of the five groups in detail below. Only four OTUs with significant abundances did not fit into any category (Table S6), including a Nitrospina relative, matching full-length sequence BS001, and a SUP05

Gammaproteobacteria matching phylotype BS077 (SUP05). These predictions are imperfect, and we acknowledge their hypothetical nature. The purpose here is to link uncultured sequence data to environmental information, and suggest testable hypotheses as to the metabolisms of these uncultured bacteria. As proof of concept, we also discuss below how this method appropriately categorized a few known and characterized organisms.

Metabolic Group I: microaerophilic bacteria

Oxygen concentrations and fluxes were highest at $\sigma_{\theta} = 15.3$, an order of magnitude lower at $\sigma_{\theta} = 15.8$, and undetectable at $\sigma_{\theta} = 16.1$ (Fig. 6a). The first depth profile category, which is defined as high relative abundance at σ_{θ} = 15.3 (39 μ M O₂) and then a steep decrease with depth, is consistent with microaerophily. Seventy-six OTUs (4192 tags) had this type of depth profile (Table S2). Gammaproteobacteria (including BS028 (Fig. 6a) from the Arctic96BD-19 cluster and BS024 of the SAR86 cluster), and Deltaproteobacteria (an unidentified SAR324 bacterium) dominated this depth profile with significant contributions from Alphaproteobacteria and Verrucomicrobia (Fig. 6a). Both aerobic heterotrophs (e.g. phylotype BS006 in the HTCC2207 cluster of oligotrophic aerobic heterotrophic isolates; Cho & Giovannoni, 2004) and aerobic autotrophs (e.g. phylotype BS003 a relative of ammonium oxidizer genus Nitrosospira; Teske et al., 1994) shared this depth profile.

Metabolic Group II: NO₃⁻ reducing bacteria

High relative abundances at $\sigma_{\theta} = 15.3$ (4.8 μ M NO₃⁻) and $\sigma_{\theta} = 15.8 \ (2.1 \ \mu M \ NO_3^{-})$ and very low abundances at $\sigma_{\theta} = 16.1 \ (0 \ \mu M \ NO_3^- and no calculated \ NO_3^- flux)$ are consistent with nitrate utilization (Fig. 6b). This is a reasonable assumption even though some nitrate reduction may be inhibited by 39 µM O2 (Oh & Silverstein, 1999) as found at $\sigma_{\theta} = 15.3$. Nitrate reduction has been measured at higher oxygen concentrations in the laboratory, presumably in anoxic micro-environments (Korner & Zumft, 1989; Oh & Silverstein, 1999). Indeed, there was a high concentration of particulate organic carbon at $\sigma_{\theta} = 15.3$ (Fig. 1), which promotes low oxygen niches. Twenty-one OTUs (4140 tags) have this type of depth profile (Table S2). Bacteria with this depth profile are dominated by Alphaproteobacteria, including an abundant SAR11 relative BS007 (Fig. 6b), with large contributions from Actinobacteria (BS079), Verrucomicrobia, and Planctomycetes (BS126). The categorization of BS079 and BS126 as nitrate reducers is at least consistent with knowledge about their most closely related cultivated isolates (Rossetti et al., 2005; Fukunaga et al., 2009).

Metabolic Group III: Mn oxidizers

A maximum in bacterial abundance at $\sigma_{\theta} = 15.8$ is consistent with manganese oxidation, because $\sigma_{\theta} = 15.8$ is the only depth sequenced for V6 tags inside the manganese oxidation zone (defined by the positive slope of particulate manganese oxides with depth; Fig. 6c). One could also imagine some nitrate reducing organisms that are sensitive to 39 μ M oxygen having this depth profile. Thirty-two OTUs (4959 tags) have this profile (Table S3). Marine Group A (phylotypes BS100 and BS110), *Actinobacteria* (phylotype BS158), and *Deltaproteobacteria* (SAR324 phylotype BS134; Fig. 6c) are among the most common taxa with this profile.

Metabolic Group IV: CH₄ oxidizers and MnO₂ reducers

As there were upward methane fluxes at $\sigma_{\theta} = 15.8$ and σ_{θ} = 16.1, V6 tag sequences of methane-oxidizing bacteria should be present at both depths. Ammonium fluxes mirrored methane fluxes (Table 1); however, organic matter respiration also produces ammonium throughout the water column. Manganese oxide reducers could also share this profile; $\sigma_{\theta} = 16.1$ is in the manganese oxide reduction zone (defined by a negative slope of particulate manganese oxides with depth; Fig. 6d) while some manganese reduction is also expected to occur at $\sigma_{\theta} = 15.8$. Methane oxidation and manganese reduction have even been linked in marine sediments (Beal et al., 2009), although whether or not they are linked in the Black Sea remains unclear. Twenty-two OTUs (3372 tags) have a Group IV depth profile (Table S4), and Gammaproteobacteria (including phylotype BS129; Fig. 6d) and Marine Group A (including phylotype BS137; Fig. 6d) were the most common taxonomic groups. Sequences affiliated with Methylococcales (type I methane oxidizers; Bowman et al., 1993) were found in this category, as were anammox bacteria related to Cand. Scalindua sorokinii. Although their V6 tag depth profiles are similar, from TRFLP, it seems likely that BS129 and BS137 (Fig. 6d) use different metabolisms. While Gammaproteobacterium BS129 smoothly increases with depth, Marine Group A BS137 has a sharp TRFLP peak maximum at $\sigma_{\theta} = 16.0$. Gammaproteobacterium BS129 is closely related to sequence II8-19 (GU108534; Fig. 3), which was enriched in ¹³C during a SIP experiment with bicarbonate (Glaubitz et al., 2010). It might be expected that a methane-oxidizer would obtain its carbon from methane rather than bicarbonate. Ammonia-oxidizers, however, are autotrophic, and a gammaproteobacterial mRNA sequence of the ammonium-oxidizing gene amoA was dominant in the lower suboxic zone in August 2005 (Lam et al., 2007). Marine

Group A BS137 seems more likely to be a manganesereducer, because its TRFLP peak height is highest at depths when net manganese oxide consumption is occurring (Fig. 6d).

Metabolic Group V: sulfur cycling bacteria

It is likely that the upward flux of sulfide affects the microbial community at $\sigma_{\theta} = 16.1$ (Table 1). Sulfide was not measurable at this depth, but the detection limit for the method used was 3 µM (Konovalov et al., 2003). As oxygen and nitrate are absent at this depth and MnO₂ and nitrite concentrations (0.04 µM) are low, fermentation and sulfate reduction may occur. Fifty-three OTUs (1178 tags) have this profile (Table S5), and the most dominant taxa are Deltaproteobacteria and Gammaproteobacteria (e.g. BS136), with important contributions from Lentisphaera and Epsilonproteobacteria. The Deltaproteobacteria are mostly represented by the group Desulfobacteraceae, which are associated with sulfate reduction (Finster et al., 1997) and fermentation (Kendall et al., 2006). Epsilonproteobacterial pyrosequences matched fulllength sequence BS139 (Fig. 6e) of the Sulfurimonas genus (Fig. 3), a genus in which all the cultured isolates mediate sulfur oxidation (Inagaki et al., 2003; Takai et al., 2006).

In summary, although we cannot attribute a metabolism to an uncultured sequence, we can make useful, testable hypotheses by examining the depth profiles of individual phylotypes across a chemical gradient. These hypotheses should focus future investigations into the physiology of specific organisms in the Black Sea.

Aggregate-attached bacteria

Both bulk water and 30 µm pore size filter samples were obtained from the center of the suboxic zone $(\sigma_{\theta} = 15.8)$. Aggregates larger than 53 µm dominate the vertical mass flux in the ocean (Clegg & Whitfield, 1990; Amiel et al., 2002). Therefore, bacteria caught on the 30 µm filter are likely to be mainly attached to sinking aggregates with some attached to large suspended particles. S-POC decreased from 14 µM in the oxic zone to 4 µM in the suboxic zone (Fig. 1) and C/N ratios were around 9 (C. Fuchsman, unpublished data), indicating that the particulate organic matter was at least partially degraded. The particulate sample is from a depth $(\sigma_{\theta} = 15.8)$ above the zone of chemosynthesis (Yılmaz et al., 2006), and so this organic material probably sank from the euphotic zone. The presence of V6 tags related to diatom chloroplasts in the 30 µm fraction (Fig. 2) is consistent with a source from the euphotic zone. Particulate manganese oxides were also present at this depth (Table 1).

These data are one of the first instances of aggregateassociated bacteria being examined under suboxic conditions. Although the Black Sea suboxic zone is only 30–40 m wide, there should still be time for the low oxygen conditions to affect the aggregates. In 1988, the average settling speed of aggregates (0.5–5.5 mm diameter) in the suboxic zone was 11.7 m day⁻¹ (Diercks & Asper, 1997). So, an average particle at $\sigma_{\theta} = 15.8$ would have been under truly suboxic conditions for 10 m, or almost 1 day. A day is long enough to allow shifts in microbial communities (McCarren *et al.*, 2010).

The taxonomic classification of aggregate-associated bacteria in the Black Sea suboxic zone was significantly different from aggregate-associated bacteria in oxic environments. In the oxic Santa Barbara Channel, Bacteroidetes, Planctomycetes, and Gammaproteobacteria dominated in sinking aggregates (DeLong et al., 1993), and in the highly oxygenated Arctic ocean, Gammaproteobacteria of the uncultured Arctic96B-1 and OM60 groups dominated aggregate-associated (> 60 μ m) clone libraries (Kellogg & Deming, 2009). In the Black Sea suboxic zone, Marine Group A, Deltaproteobacteria, and Planctomycetes dominated the aggregate-associated fraction (> 30 µm) (Fig. 2). Many other groups were enriched in the Black Sea aggregate-associated fraction, including Lentisphaera, WS3, and Epsilonproteobacteria. The aggregate-associated community in the Black Sea suboxic zone appears to be distinct from those found in oxic environments.

We consider the aggregate-associated community to be the bacteria on the 30 µm filter. However, some abundant free-living bacteria may have been trapped in the 30 µm filter. Therefore, we have distinguished between aggregate-associated bacteria and bacteria expected to be actually aggregate-attached according to their relative abundance in the bulk water sample compared with the 30 µm filter. We consider aggregate-attached bacteria to have $5 \times$ greater abundance in the particulate sample than in the bulk water sample, and consider free-living bacteria to have $4 \times$ greater abundance in the bulk water. In this case, even though SAR11 relatives were found on the 30 µm filter, they are considered free-living due to their extremely high abundance in the bulk water sample. Of the 125 OTUs with more than five V6 tag sequences found in the 30 µm sample, 53 OTUs are considered here to be aggregate-attached (Fig. 5c, column 1 and Table S7). Thirty-seven OTUs, with more than five V6 tag sequences found in the σ_{θ} = 15.8 water sample, had 4× greater abundance in the bulk water, and are considered here to be free-living (Fig. 5c, column 2), whereas 36 OTUs were abundant in both the aggregate and bulk water samples at $\sigma_{\theta} = 15.8$ (Fig. 5c, column 3). Free-living and aggregate-attached OTUs have strikingly different community structures (Fig. 5c). Alphaproteobacteria in

the SAR11 clade dominated the free-living fraction. *Planctomycetes*, *Deltaproteobacteria*, Marine Group A, *Lentisphaera*, *Epsilonproteobacteria*, WS3, and *Deinococci* were enriched in the aggregate-attached fraction. A large number of OTUs (111) were shared only between the aggregate-attached fraction at $\sigma_{\theta} = 15.8$ and bulk water at $\sigma_{\theta} = 16.1$, a depth influenced by sulfide (Fig. 5b). These OTUs are shared in spite of the fact that the sinking aggregates would not yet have had contact with that depth. These OTUs include members of the *Epsilonproteobacteria*, WS3, *Lentisphaera*, and *Deinococci*. Some OTUs found in the aggregate-attached fraction were not significant in the bulk water at any depth. These include some members of the *Planctomycetes* (BS097, BS109) and the *Deltaproteobacteria*.

The overlap of many OTUs between the particulate sample at σ_{θ} = 15.8 and the bulk water sample at σ_{θ} = 16.1, a sample that was influenced by the sulfide flux from the anoxic zone (Table 1), suggests that S cycling may be occurring in the aggregate. This is supported by the presence of aggregate-attached Epsilonproteobacteria from the Sulfurimonas genus (BS139) closely related to S. denitrificans and S. autotrophica (Fig. 3). All cultured members of this genus have been found to oxidize sulfide with NO₃⁻ or O₂ (Inagaki et al., 2003; Takai et al., 2006). Clone BS077, a SUP05 which is aggregate-associated and also abundant in the bulk water sample, is closely related to SUP05 from Saanich Inlet (Fig. 3), which is implicated in sulfur oxidation by metagenomic sequences (Walsh et al., 2009). Sulfur oxidation genes from SUP05 were also present and expressed in an oceanic oxygen minimum zone (Canfield et al., 2010; Stewart et al., 2011). Many aggregate-attached V6 tags were also assigned to Desulfobulbacaea and Desulfuromonadales, two potentially sulfate reducing orders of bacteria. Thus, we have candidates for both sulfate reduction and sulfide oxidation potentially attached to particles. In addition, aggregate-attached Marine Group A bacterium BS137 is associated with manganese reduction (Fig. 6e). Aggregate-associated bacteria also include Marine Group A bacteria, which are linked to manganese oxidation (e.g. BS110), and Actinobacteria (e.g. BS079) associated with nitrate reduction (Fig. 6; Tables S3 and S4). Therefore, we predict that the aggregates hosted a variety of metabolisms.

It is unclear whether sulfate reduction is feasible at $\sigma_0 = 15.8$, because sulfate reduction is typically inhibted by oxygen. Sulfate reduction has been measured in the Oxygen Minimum Zone off of Chile (Canfield *et al.*, 2010) even though sulfide concentrations were below detection. Oxygen concentrations in the OMZ, however, were significantly lower than those measured at $\sigma_0 = 15.8$ in the Black Sea (20 nM vs. 2 μ M). In the following, we evaluate the possibility of sulfur reduction inside aggregates from the suboxic zone with a simple model of oxy-

gen penetration into the aggregate. Nutrient gradients in and around aggregates can be described using molecular diffusion across a diffusive boundary layer that surrounds the aggregates (Alldredge & Cohen, 1987; Ploug et al., 1997). Oxygen utilization in particles is not transportlimited in fully oxygenated seawater, but rather reactionlimited, which is why aggregates are generally not anoxic (Ploug, 2001). However, calculations predict that at c. 25 µM ambient O2, oxygen utilization becomes transport limited (Ploug, 2001). Both oxygen (2 µM) and nitrate (2 μ M) should be transport limited at $\sigma_{\theta} = 15.8$. The volumetric oxygen respiration rate for a 1 mm diameter particle with a 0.17 mm boundary layer thickness (after Ploug et al., 1997) is 162 µmol cm⁻³ day⁻¹ organic carbon in the aggregate at $\sigma_{\theta} = 15.3$, but much lower (4 μ mol cm⁻³ day⁻¹) in the aggregate at $\sigma_{\theta} = 15.8$. It seems reasonable to expect that oxygen would not reach the center of such an aggregate at $\sigma_{\theta} = 15.8$.

We suggest that an aggregate under suboxic conditions contains multiple niches. An outer layer would contain organisms that utilize nitrate or oxygen, whereas in the center of the aggregate, sulfate reduction and manganese reduction might occur. Sulfate reduction produces sulfide (and ammonium) that would be oxidized in the outer layer of the particle. This layering of aerobic and anaerobic bacteria in aggregates under suboxic conditions has been seen in wastewater treatment plants (Vlaeminck et al., 2010). Thus, we expect a large diversity of metabolisms to occur in a relatively small volume. This diversity of metabolisms under low oxygen conditions may be why the V6 region of 16S rRNA gene aggregate-associated community was the most diverse in this study (Fig. 5a). Our results contrast to the highly oxygenated Arctic Ocean, where the aggregateassociated bacterial community was less diverse than the free-living community (Kellogg & Deming, 2009). This difference is consistent with the fact that in oxygen-saturated water, aggregates can only undergo transitory anoxic conditions and that sustained sulfate reduction and methanogenesis cannot occur (Ploug et al., 1997).

Conclusions

The combination of in-depth sequencing of the V6 region of 16S rRNA gene to provide information about the entire microbial community, TRFLP to give spatial resolution, full-length clones to give these techniques taxonomic resolution, and extensive chemical profiles to provide ecological context have allowed us to examine bacterial communities of the suboxic and hypoxic zones of the Black Sea more comprehensively than ever before.

Five general depth profiles of bacterial abundance were identified and correlated with geochemical data, and the metabolisms that might correspond to each of these depth profiles were predicted. These predictions, although imperfect, provide testable hypotheses regarding the metabolic strategies of uncultured bacteria in the Black Sea suboxic zone. A series of metagenomic analyses of the Black Sea suboxic zone would better link bacterial identity to metabolic genes. Future work could include experiments (e.g. stable isotope probing) designed to explicitly test for predicted metabolic reactions.

Our study also highlights the effect of low redox conditions on the microbial diversity of sinking aggregates. Free-living and aggregate-attached OTUs had strikingly different taxonomies. Aggregate-attached OTUs included bacteria linked to sulfate reduction and sulfide oxidation, implying more reducing E_h conditions in aggregate interiors than found in the ambient water.

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Supporting Information

Additional Supporting Information may be found in the online version of this article:

Fig. S1. Depth profiles, using a density scale, for oxygen (squares), sulfide (triangles), nitrate (circles), nitrite (diamonds), ammonium (x), methane (crosses), and particulate manganese (bold squares) from the Western Central Gyre of the Black Sea in March 2005.

Fig. S2. Cluster of community similarities calculated from TRFLP profiles from depths throughout the oxygenated, suboxic, and sulfidic layers.

Fig. S3. TRFLP chromatograms from the suboxic zone, obtained using the enzyme MspI.

Fig. S4. TRFLP chromatograms from the deep suboxic and sulfidic zones, obtained using the enzyme MspI.

Table S1. Cell counts from the suboxic zone.

Table S2. Depth Profile Metabolic Group I: The number of normalized tags at each depth for each unique OTU in this depth profile as well as their lifestyle and taxonomy.

Table S3. Depth Profile Metabolic Group II: The number of normalized tags at each depth for each unique OTU in this depth profile as well as their lifestyle and taxonomy.

Table S4. Depth Profile Metabolic Group III: The number of normalized tags at each depth for each unique OTU in this depth profile as well as their lifestyle and taxonomy.

Table S5. Depth Profile Metabolic Group IV: The number of normalized tags at each depth for each unique OTU in this depth profile as well as their lifestyle and taxonomy.

Table S6. Depth Profile Metabolic Group V: The number of normalized tags at each depth for each unique OTU in this depth profile as well as their lifestyle and taxonomy.

Table S7. Unassigned Depth Profile: The number of normalized tags at each depth for each unique OTU in this depth profile as well as their lifestyle and taxonomy.

Table S8. Aggregate-attached bacteria: The number of normalized tags at each depth for each unique OTU as well as their taxonomy.

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